

The Classic Problem of Indian Monsoon-what AMSR-2 can reveal

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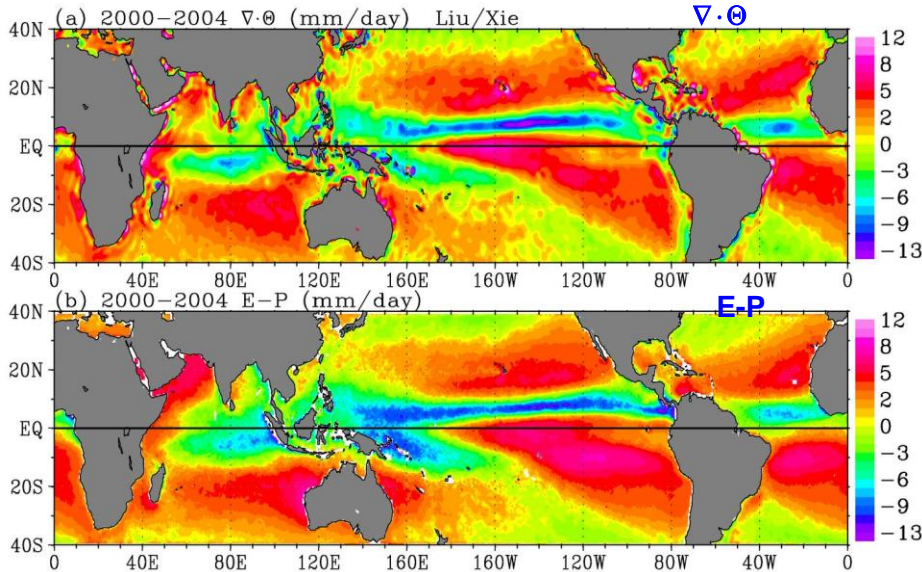
GCOM-W in Tokyo 1/15/2015

- EXTEND THE TIME SERIES WE PRODUCED WITH AMS-E DATA
- CONTINUE TO FIND NEW WAYS OF APPLYING AMSR DATA

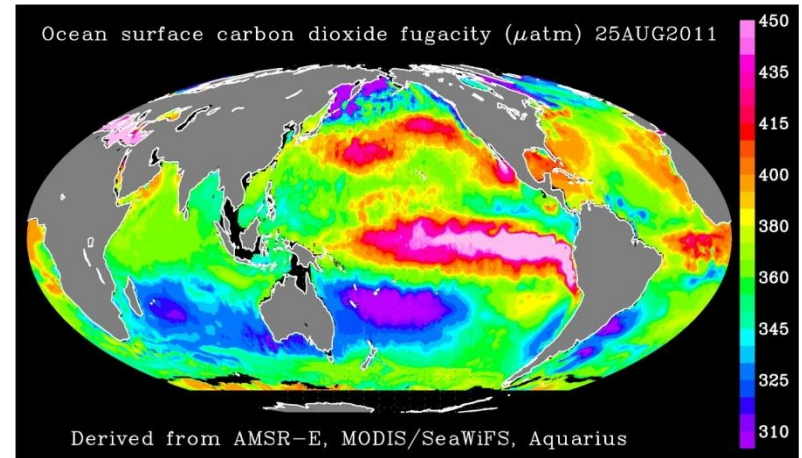
Ocean's Role in Water/Energy/Carbon Cycles

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Water Cycle

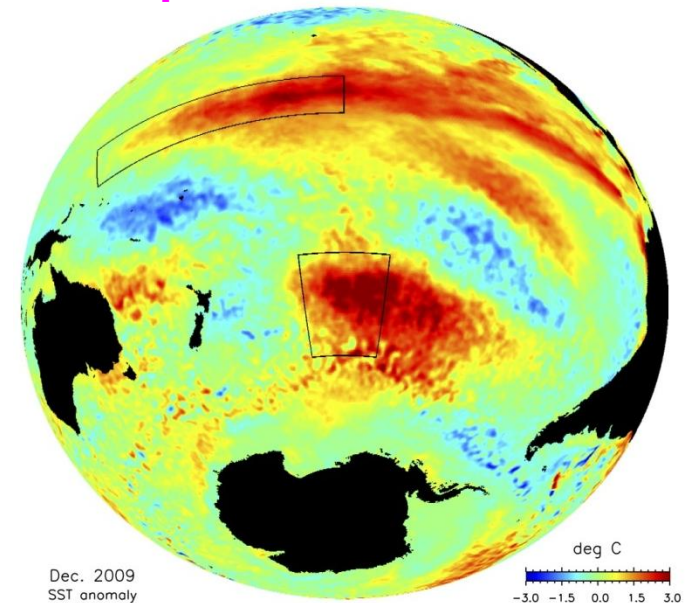
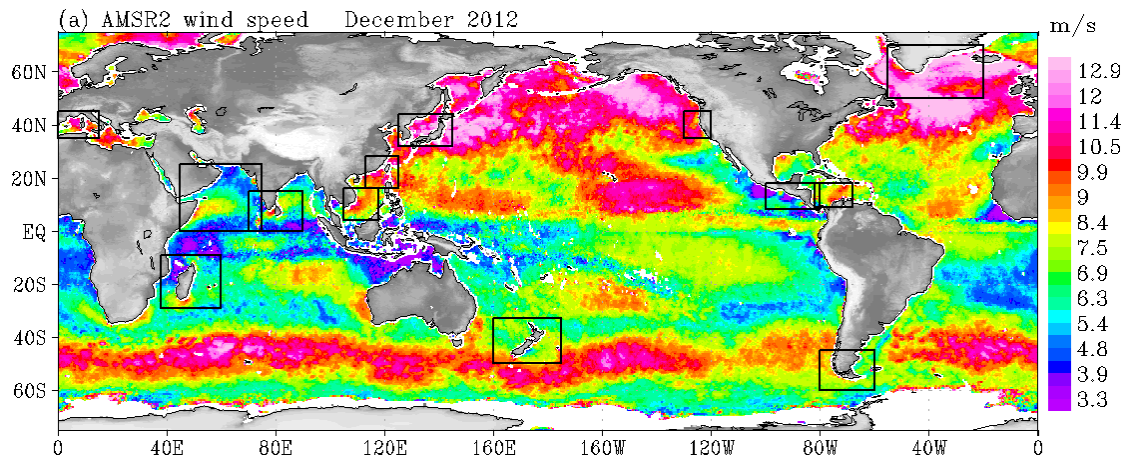


Carbon Cycle



Temperature Anomalies

Wind Farm Deployment



Extending Data Sets Derived from AMSR-E with AMSR-2

Using AMSR data, we developed and validated retrieval methods for

- (1) Ocean surface evaporation (E) through bulk parameterization**
- (2) E directly from brightness temperatures measured by AMSR-E**
- (3) Integrated water transport. The divergence of the transport is E-P**
- (4) Wind power distribution for deployment of floating wind farms**
- (5) Ocean surface carbon dioxide fugacity**

**Produced and provided accessibility to 9 years (6/2002-9/2011) data of
(1)-(5)**

<http://airsea.jpl.nasa.gov/seaflux/seaflux.html>

Monsoon-the basic definition

- **Monsoon is the seasonal reversal of transport between land and ocean that has strong rain effect**
- **Synoptic rain events should be linked to seasonal oceanic moisture supply.**
- **Interannual and decadal modifications of monsoon in terms of global teleconnection between rain and SST must be interpreted through regional ocean-land exchanges.**
- **Besides rain, soil moisture has given us an additional spacebased observation on terrestrial response to ocean influence**

HYDROLOGIC BALANCE

$$\frac{\partial W}{\partial t} + \nabla \cdot \Theta = E - P$$

$$\Theta = \frac{1}{g} \int_0^{p_0} q U dp$$

$$W = \frac{1}{g} \int_0^{p_0} q dp$$

$$\Theta = U_e W$$

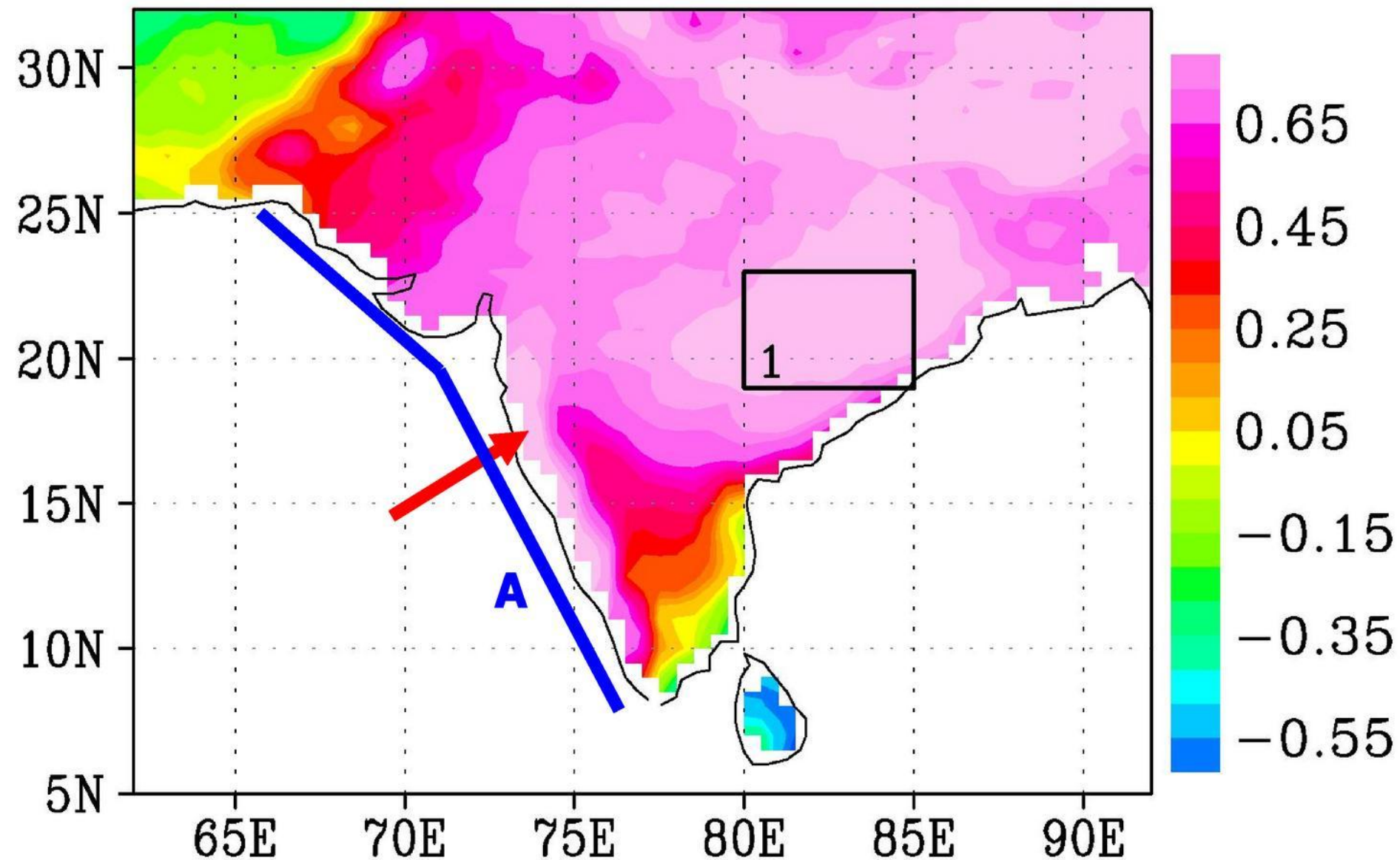
Θ is equivalent to column water vapor W advected by U_e .

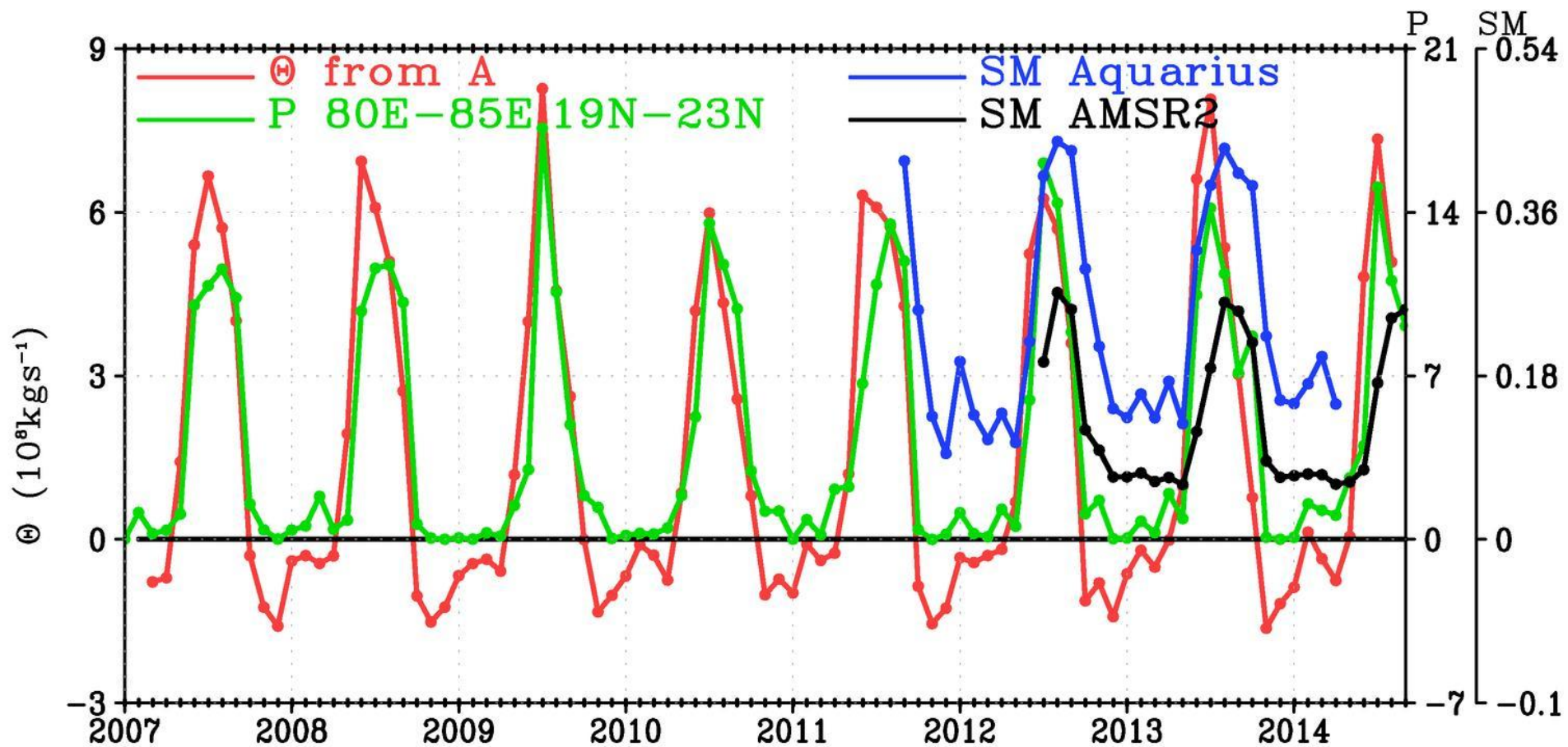
U_e is the depth-averaged wind weighted by humidity

We use SVR to relate U_e to wind at two levels:

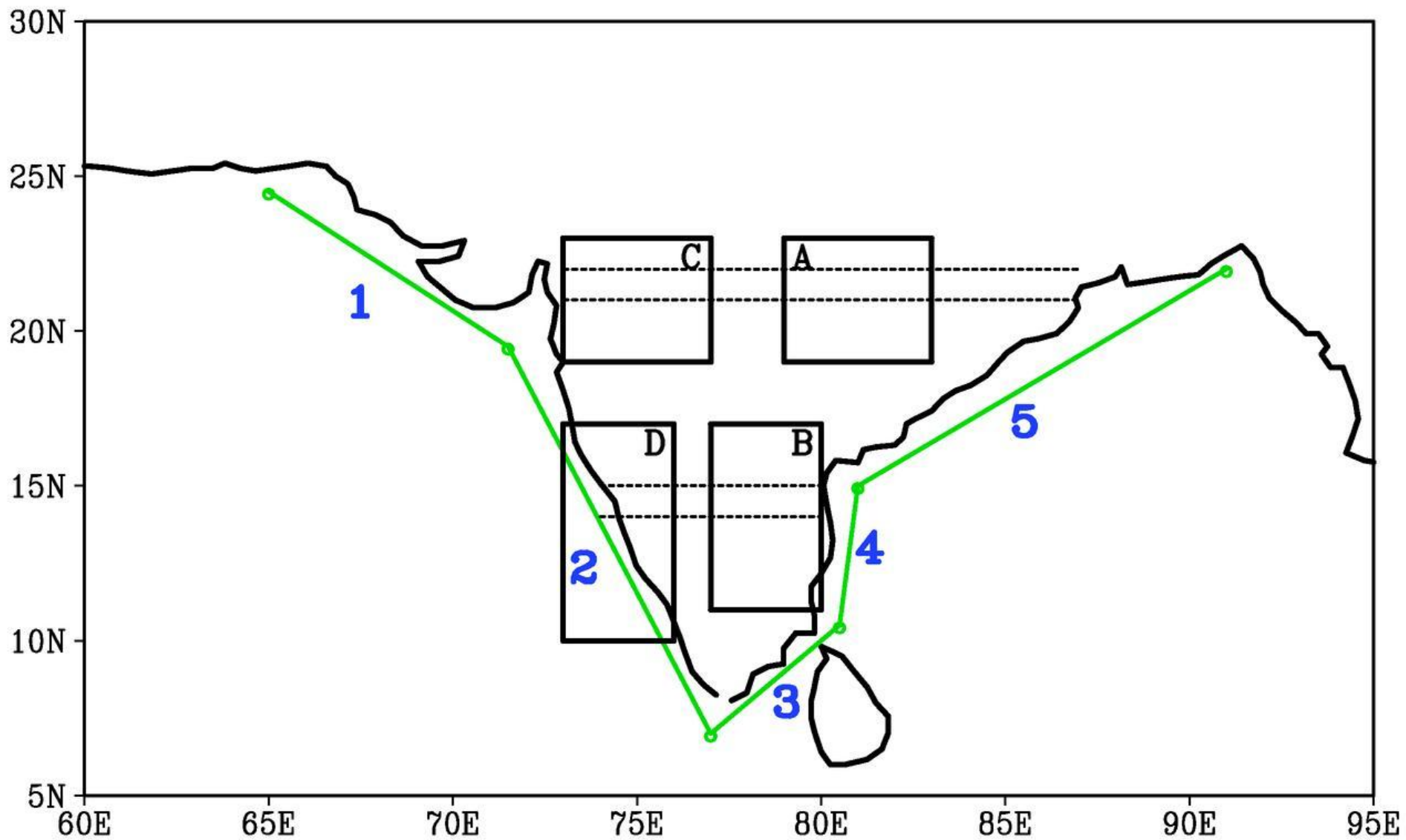
1. U_N : scatterometer surface wind stress
2. U_{850mb} : cloud drift wind (free-stream wind)

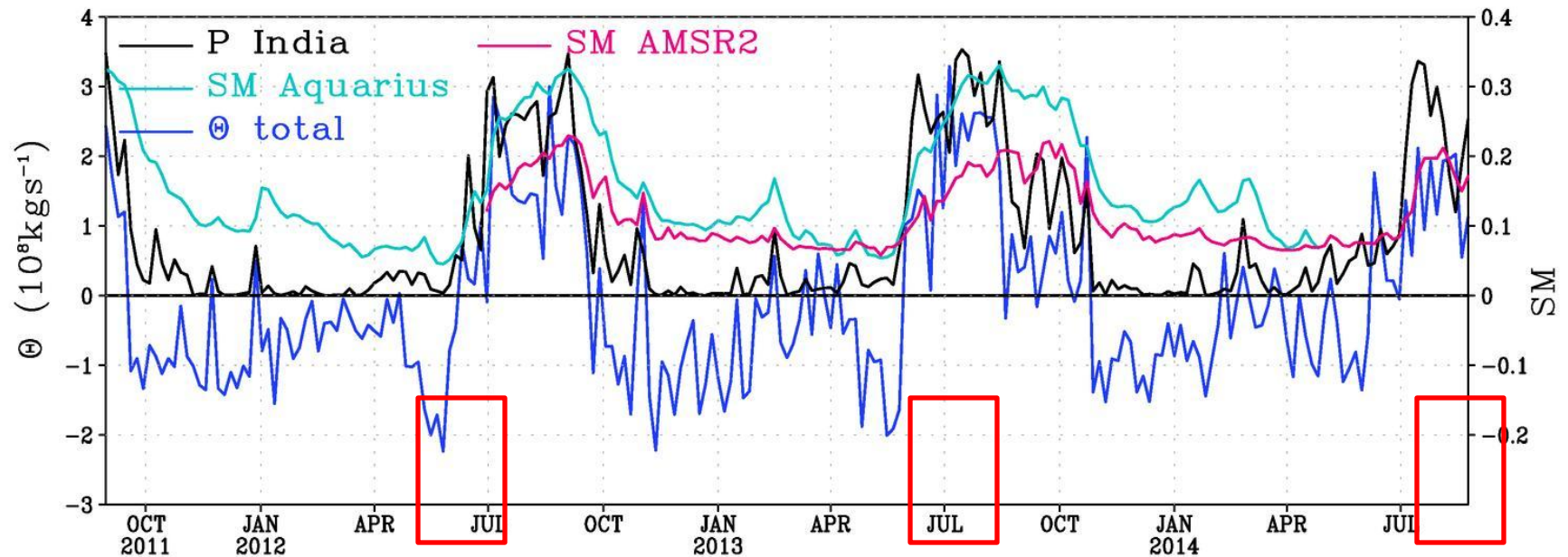
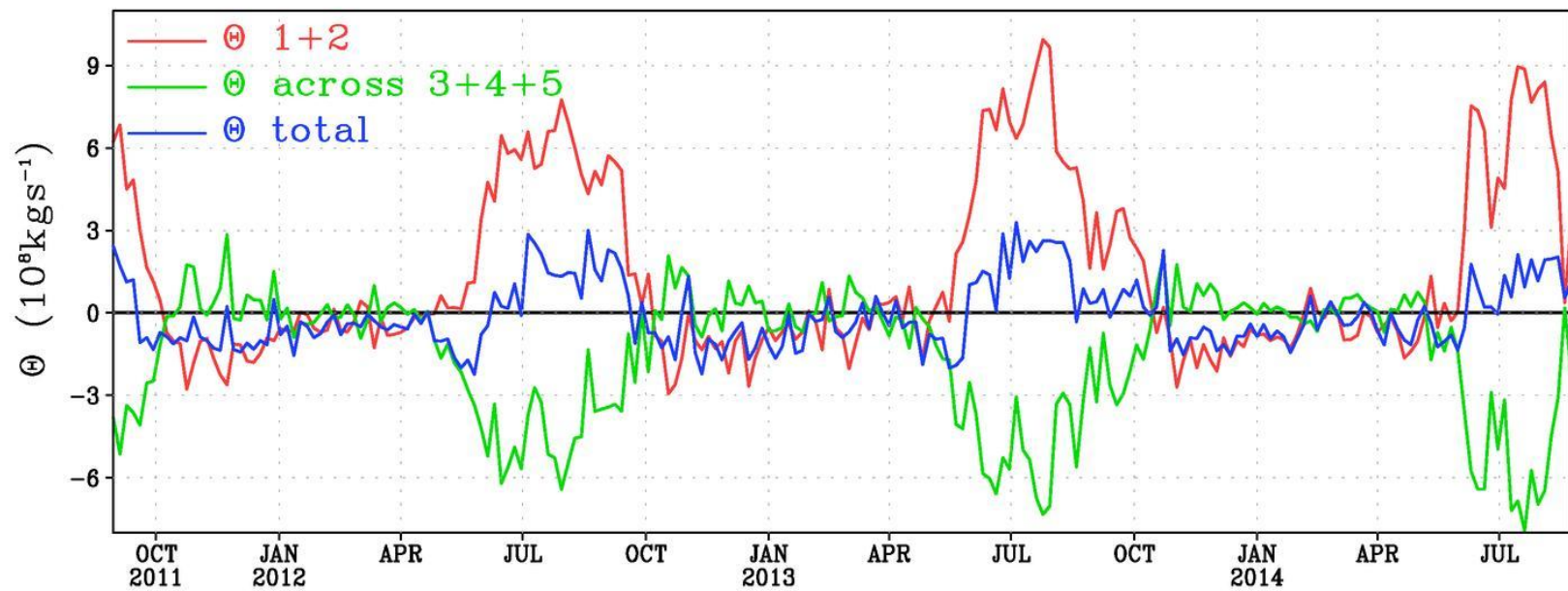
Correlation coefficient between Θ from A and precipitation



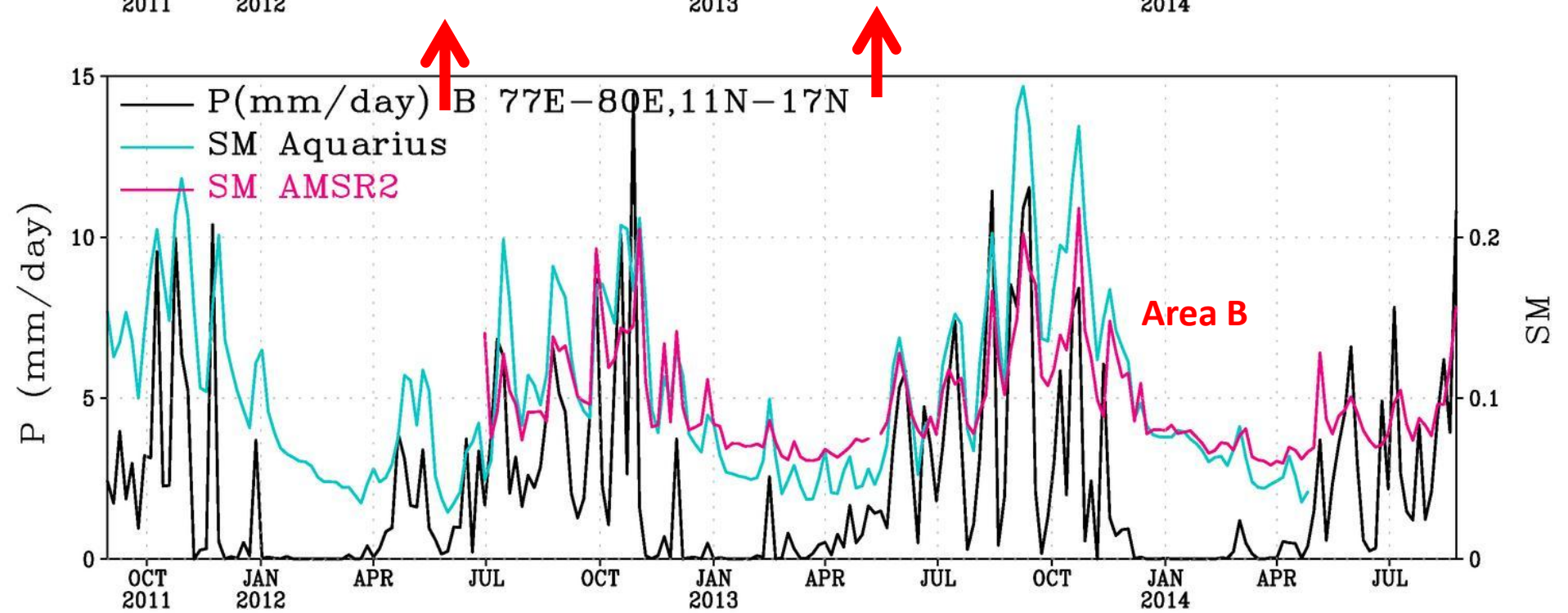
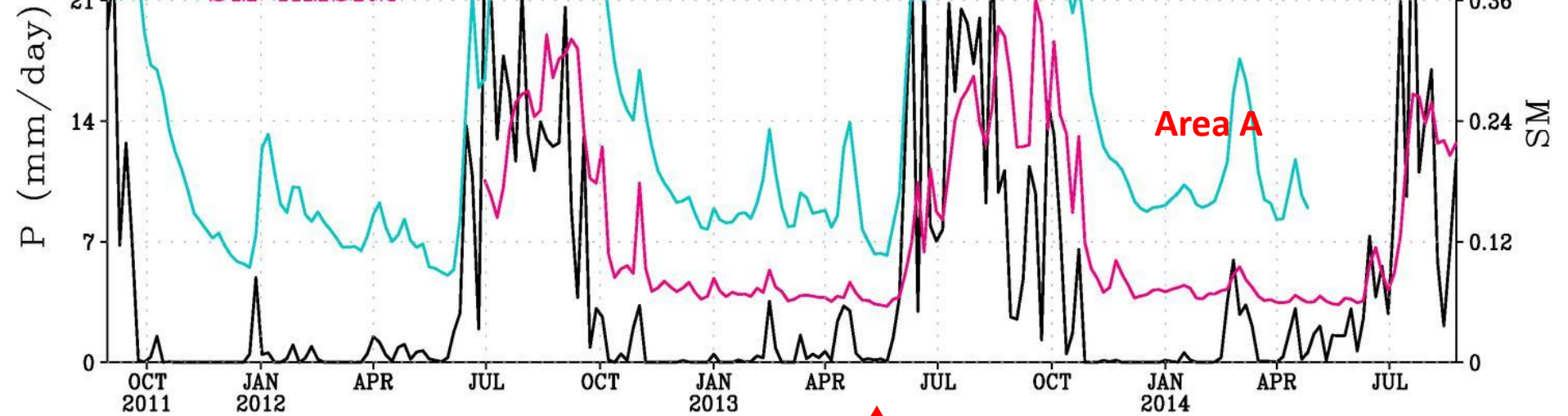


Obvious bias in soil moisture between AMSR-2 and Aquarius

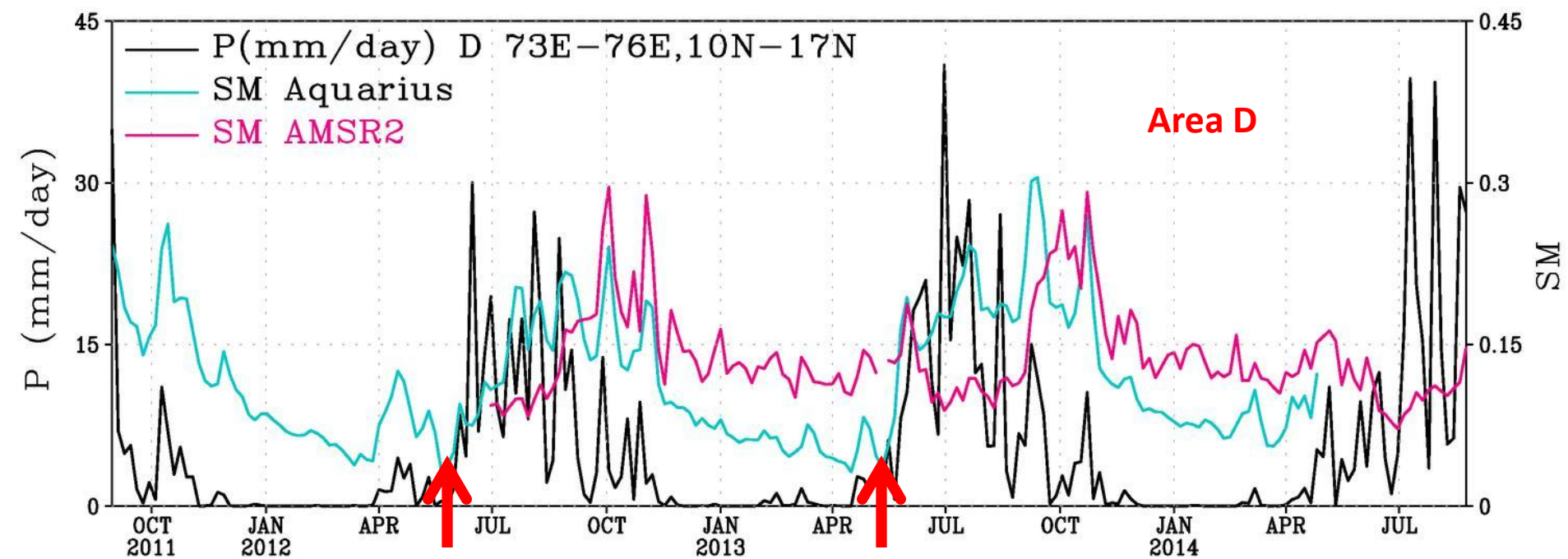
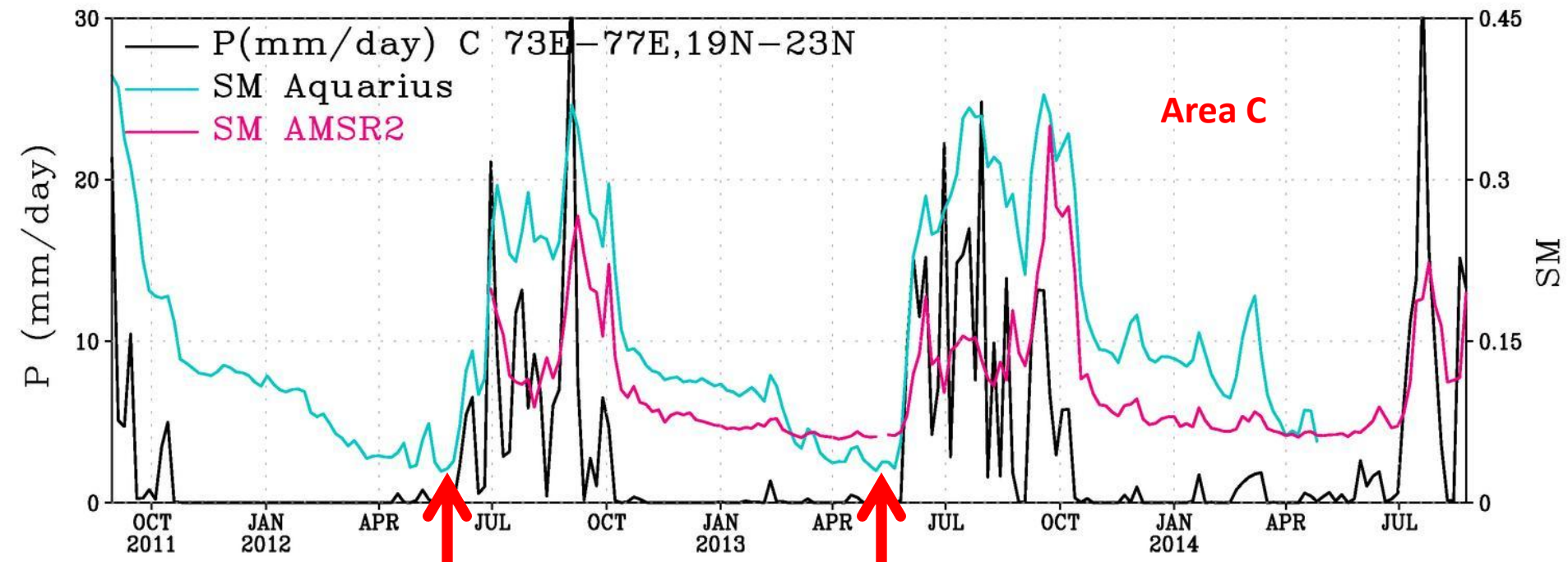


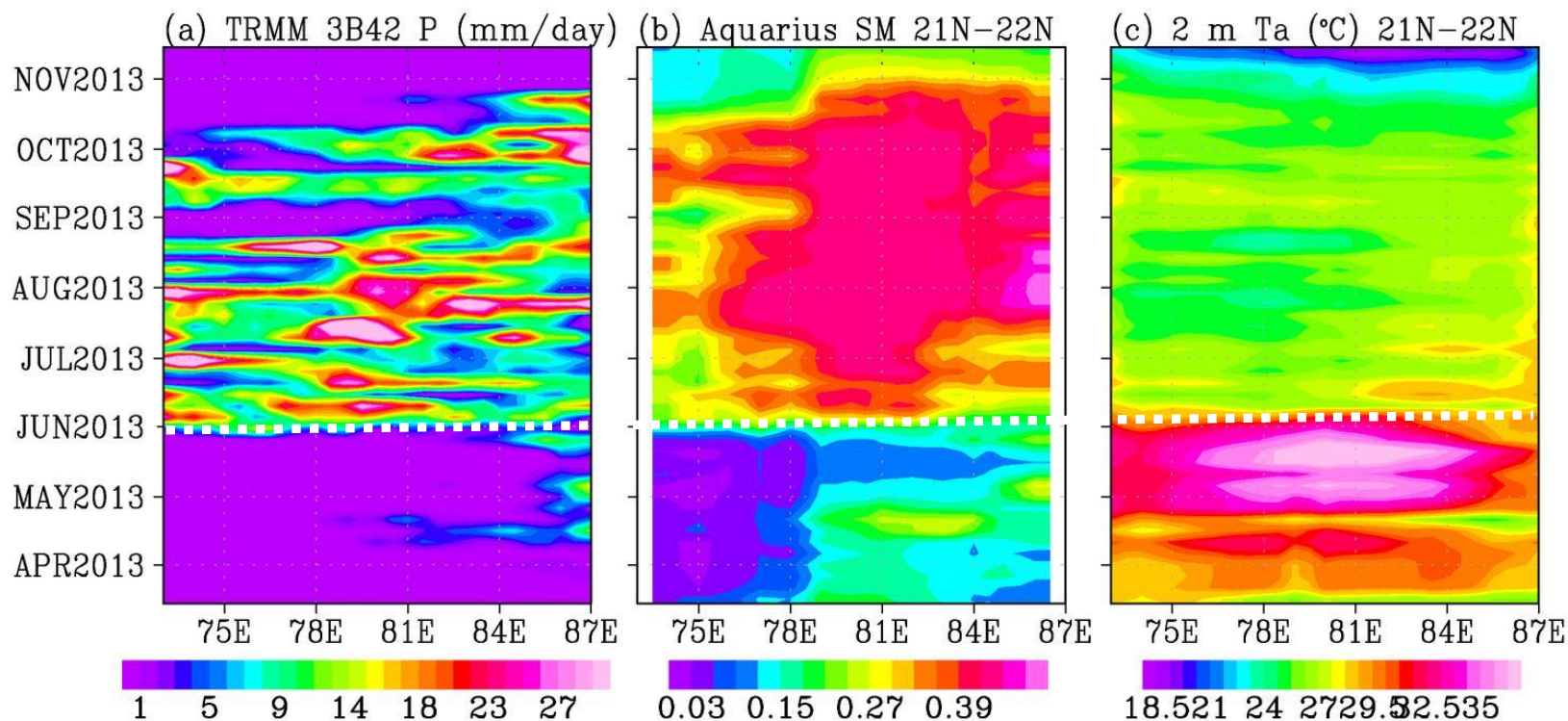
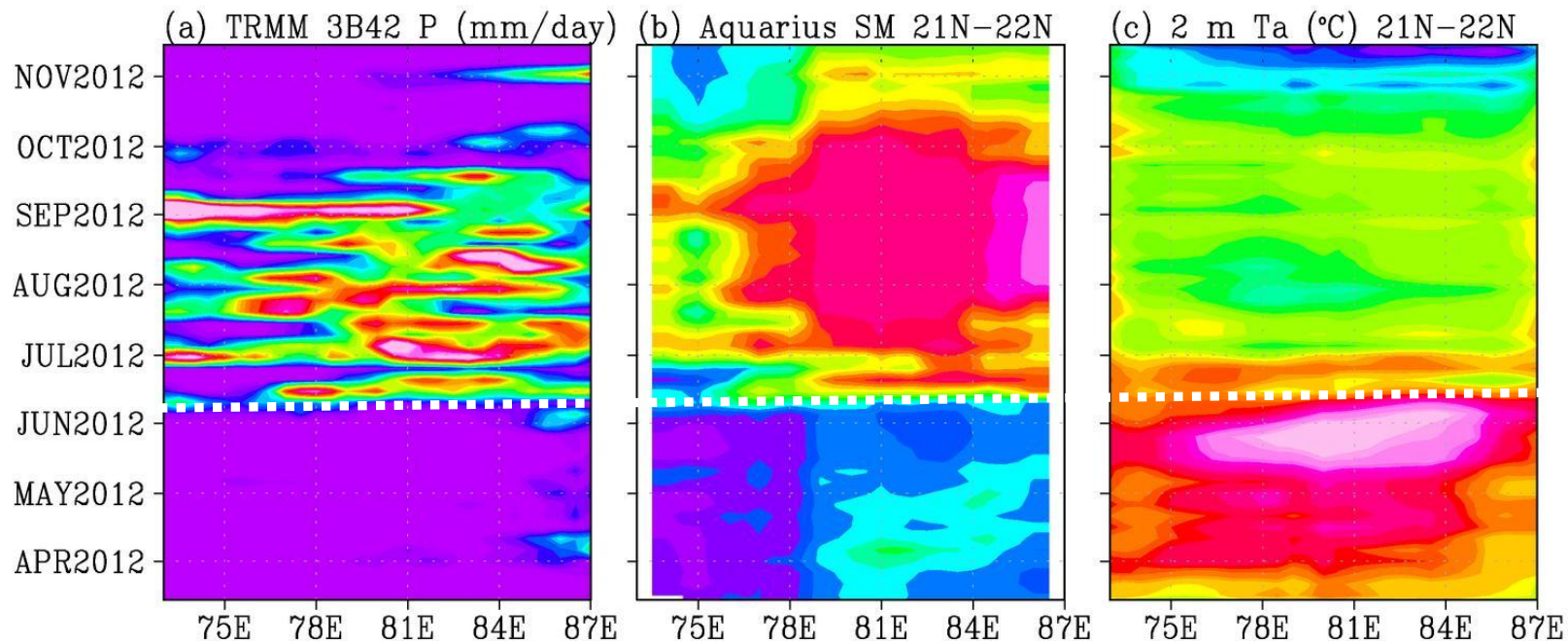


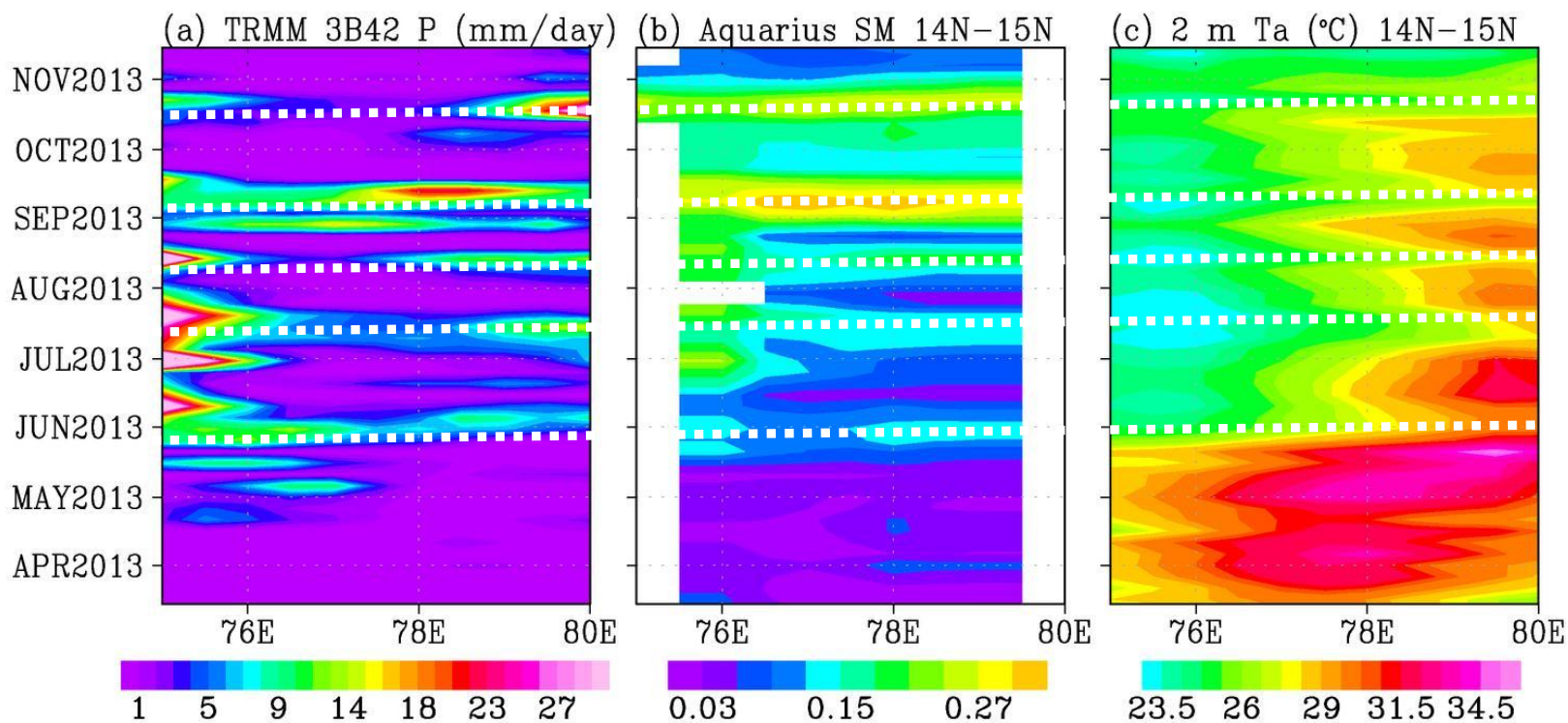
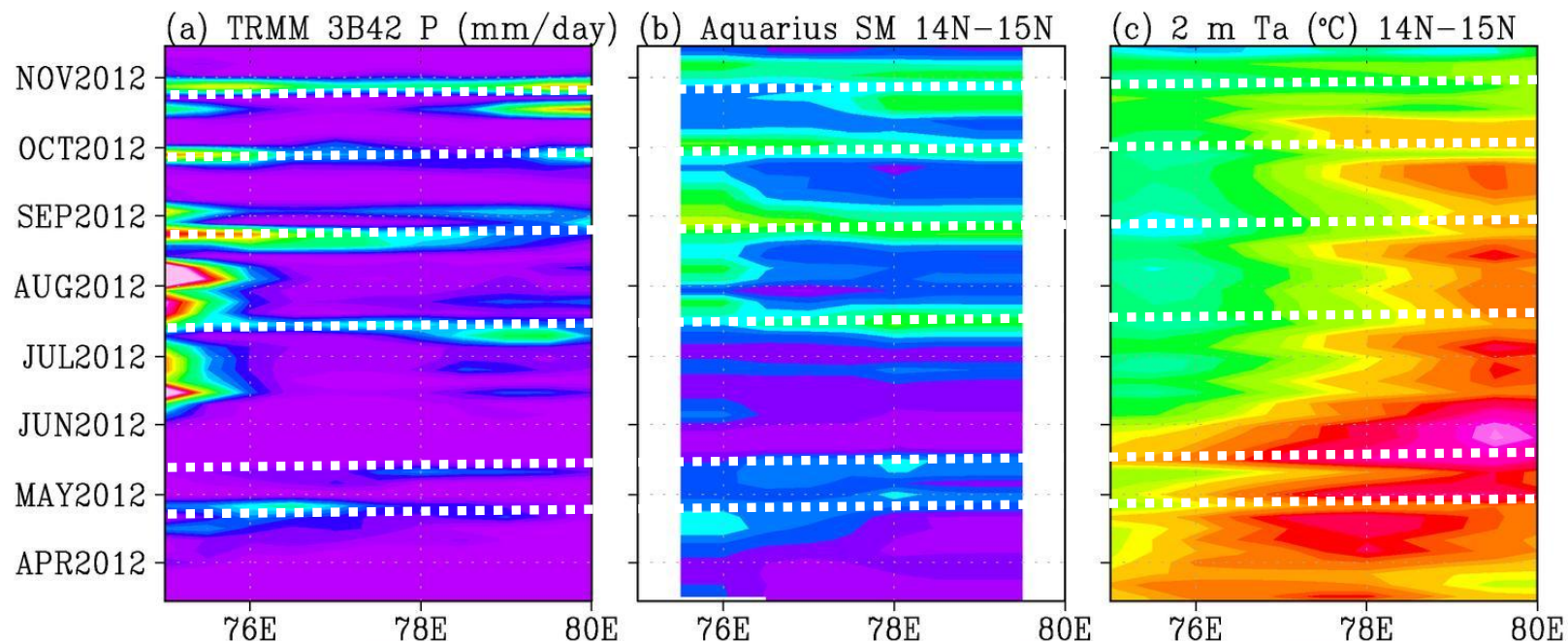
More moisture advected out to BB than coming in from AS before monsoon onsets



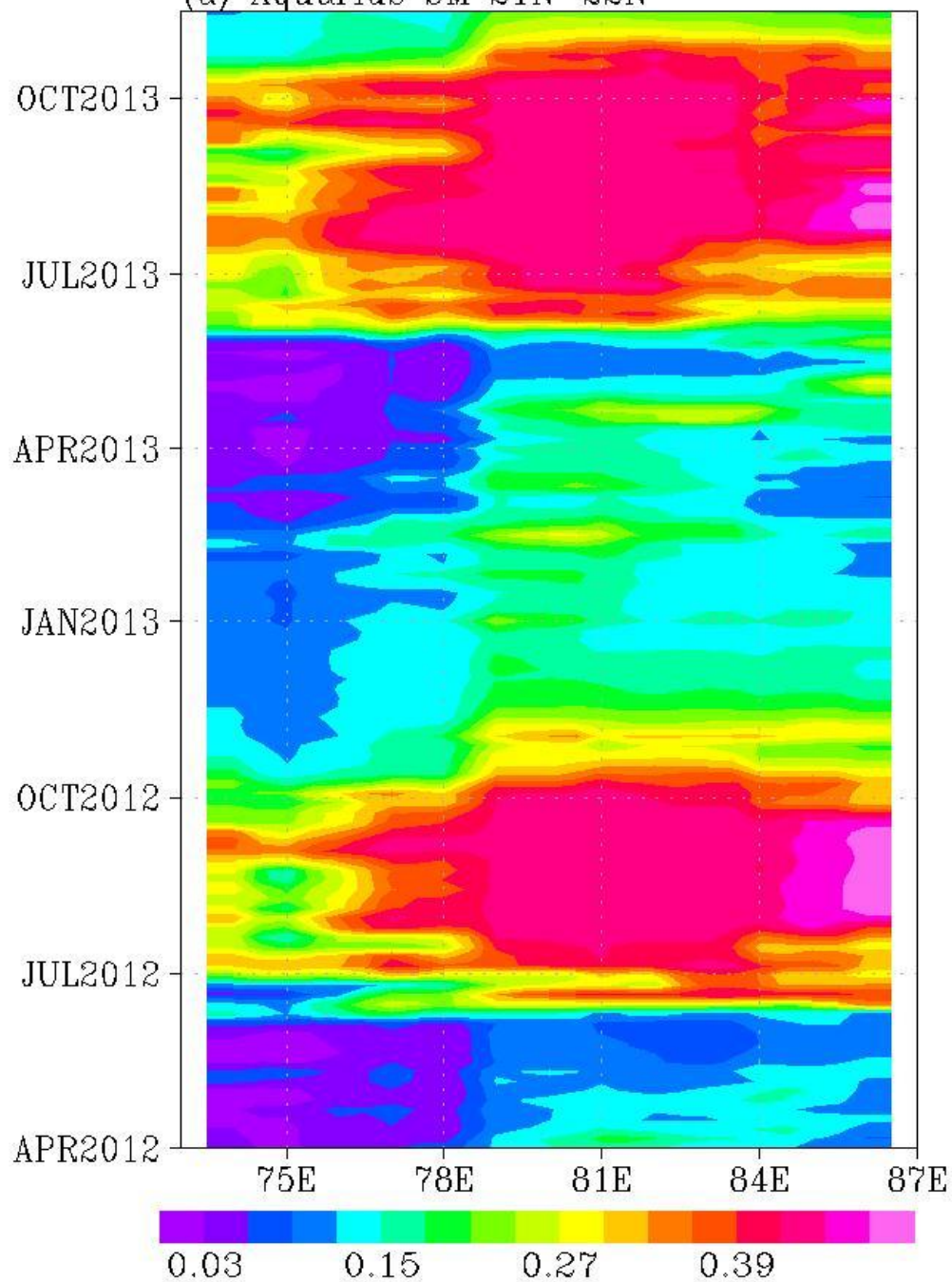
Pre-monsoon drought reflected in soil moisture



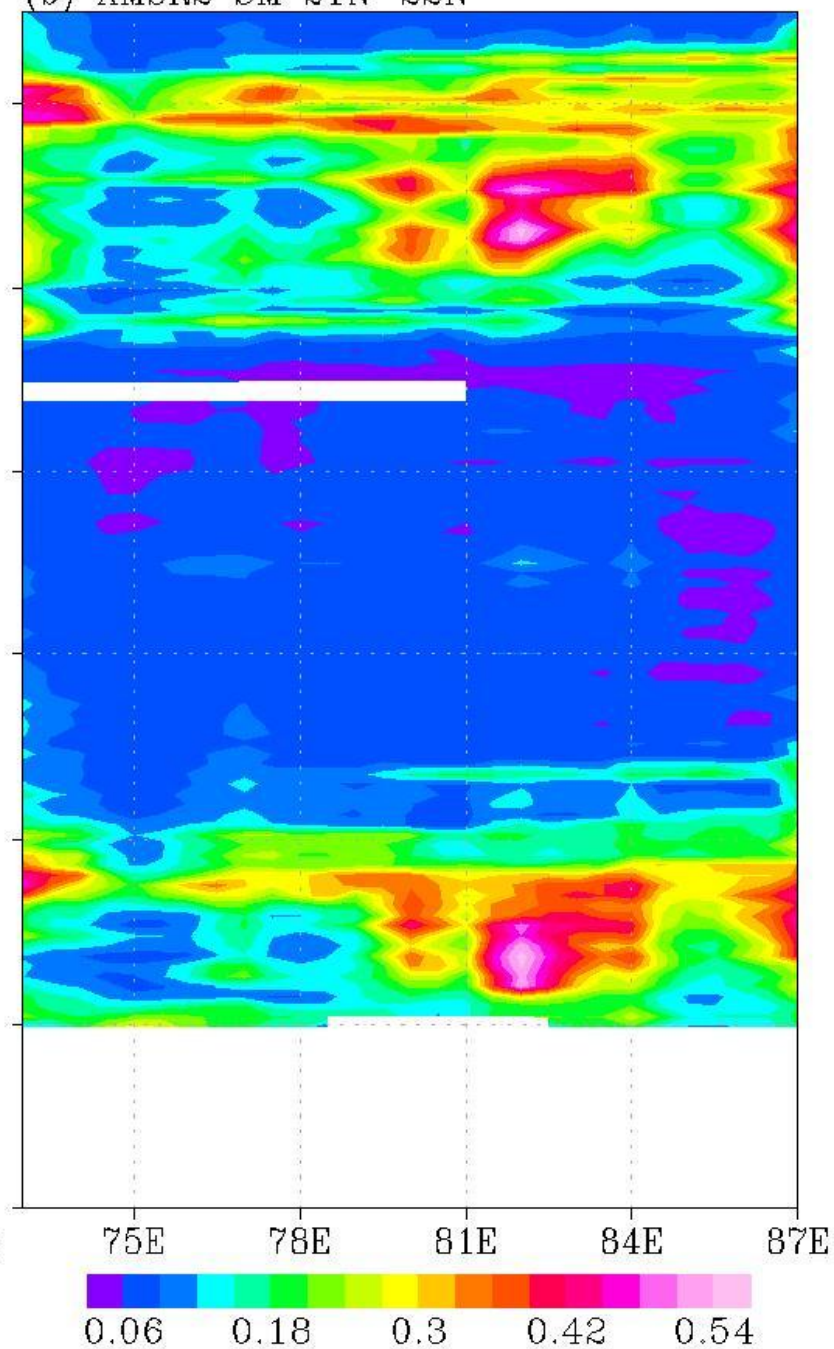




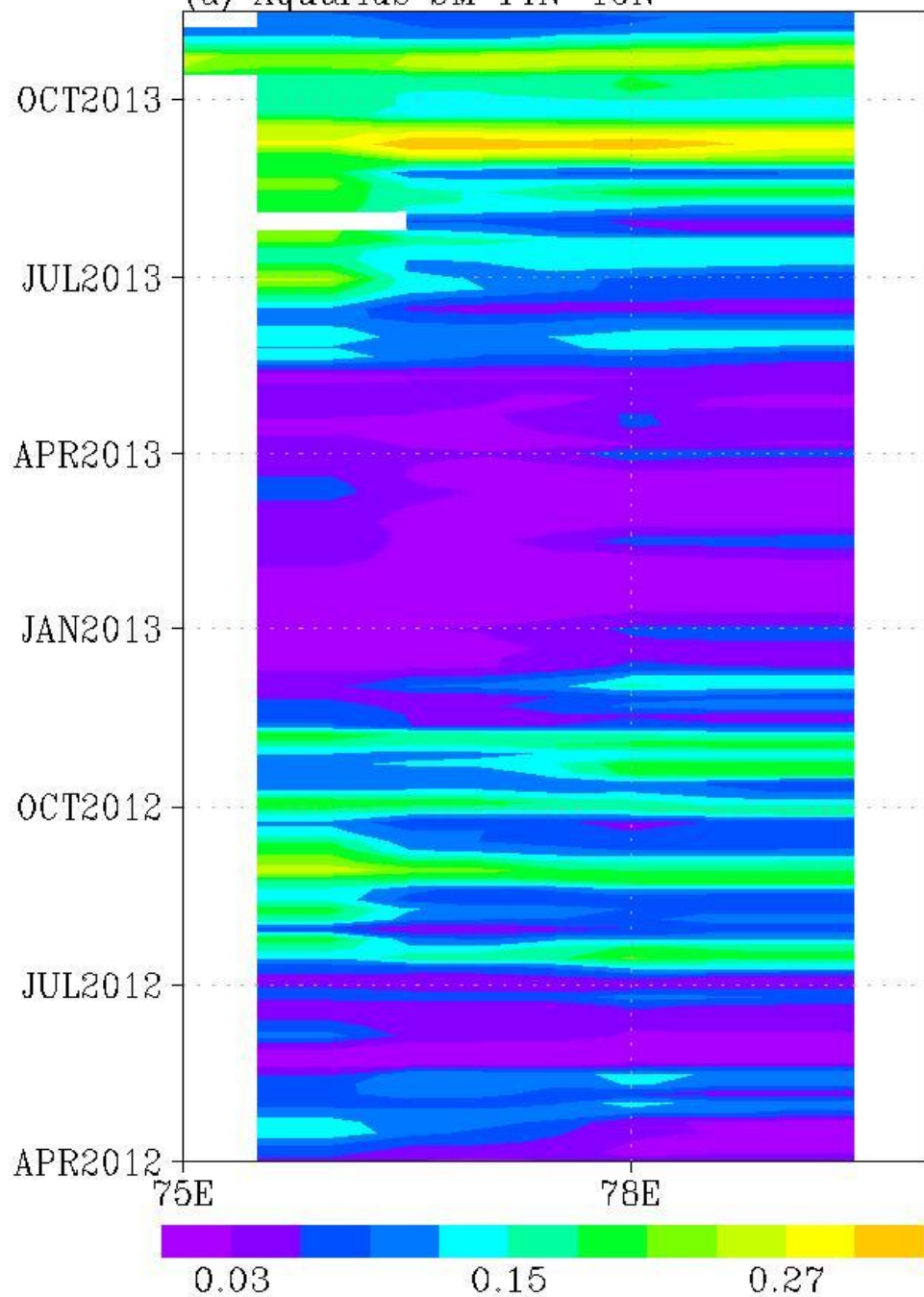
(a) Aquarius SM 21N-22N



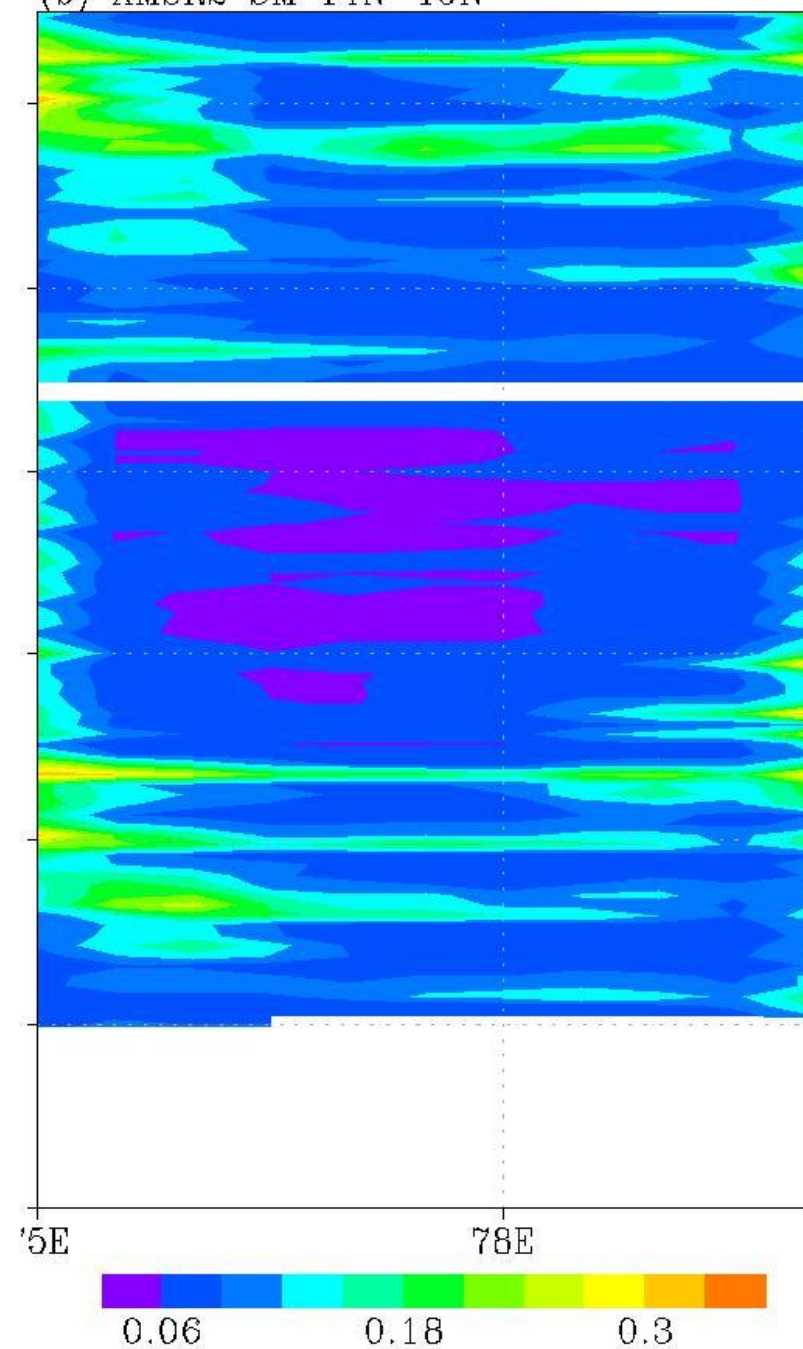
(b) AMSR2 SM 21N-22N



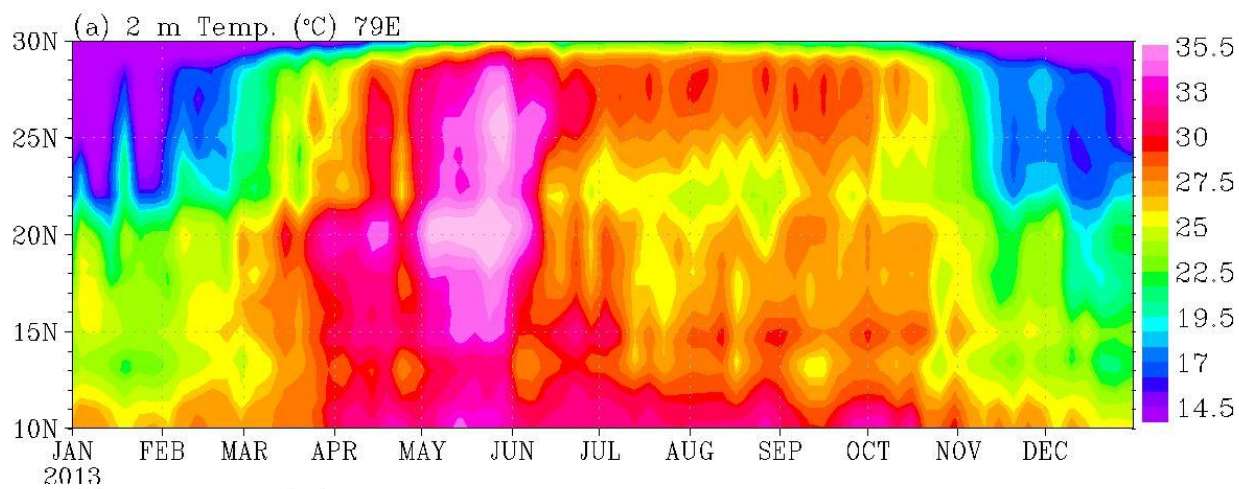
(a) Aquarius SM 14N–15N



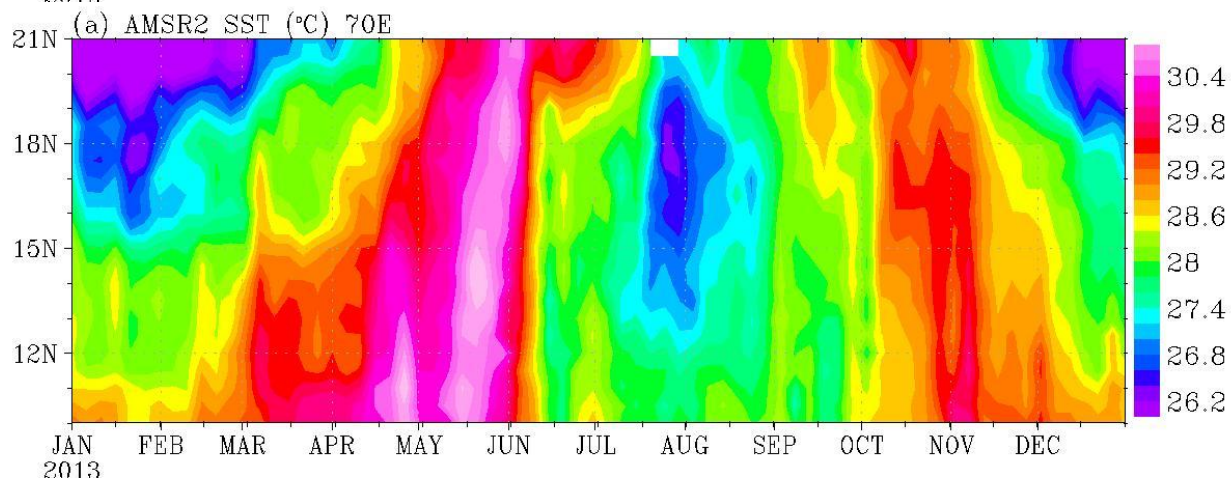
(b) AMSR2 SM 14N–15N



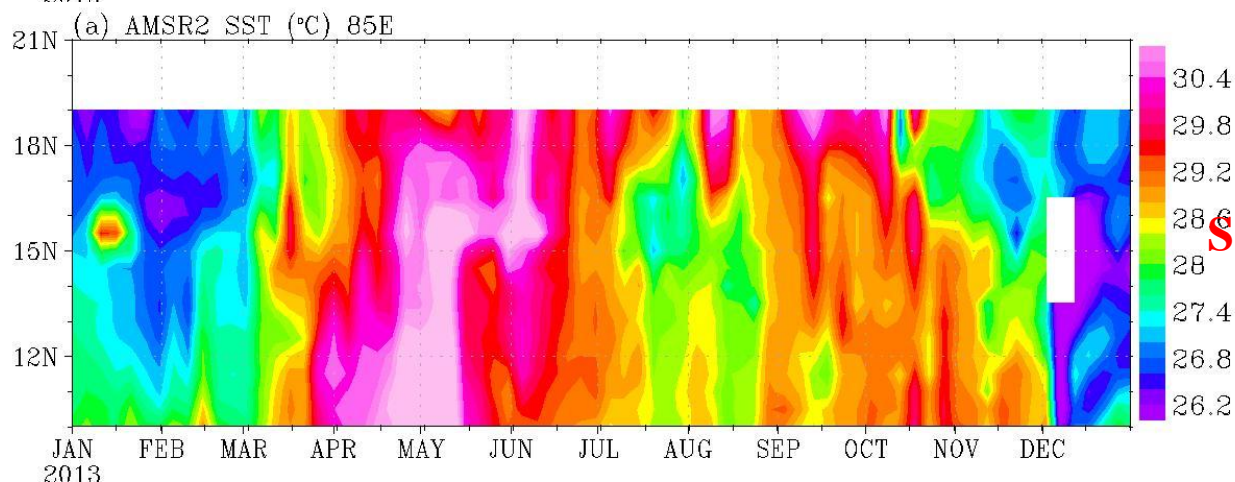
Temperature gradient



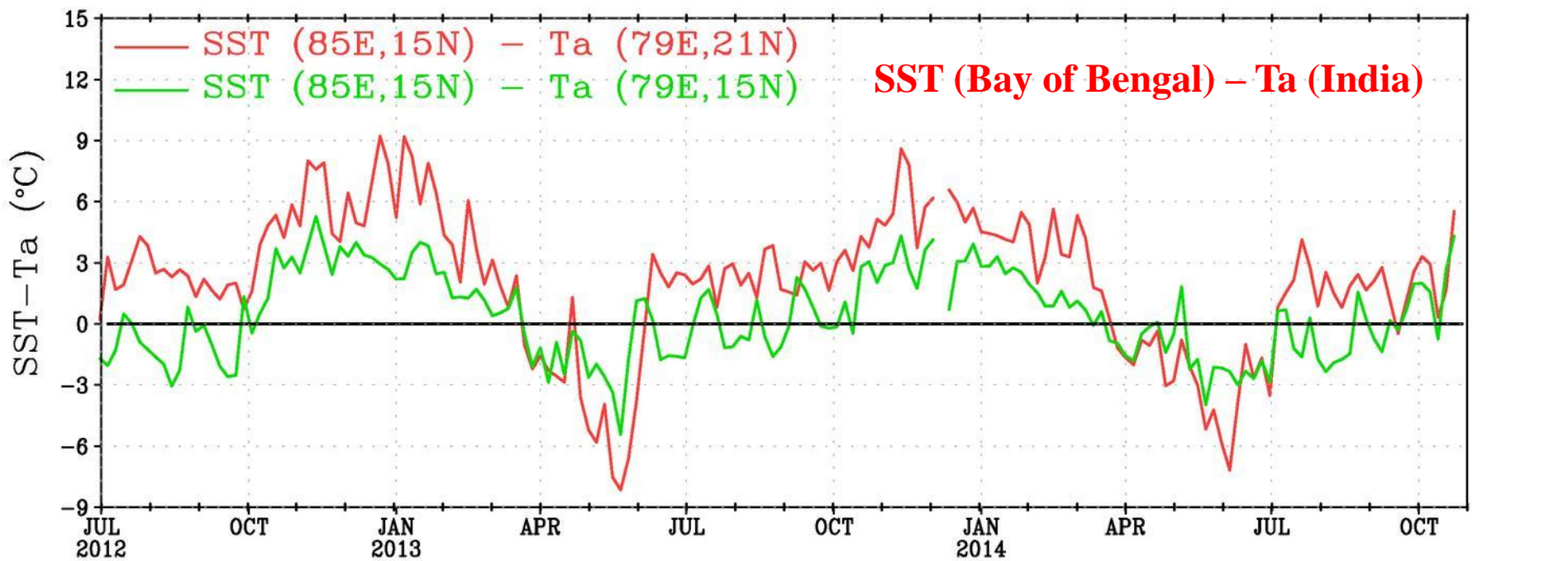
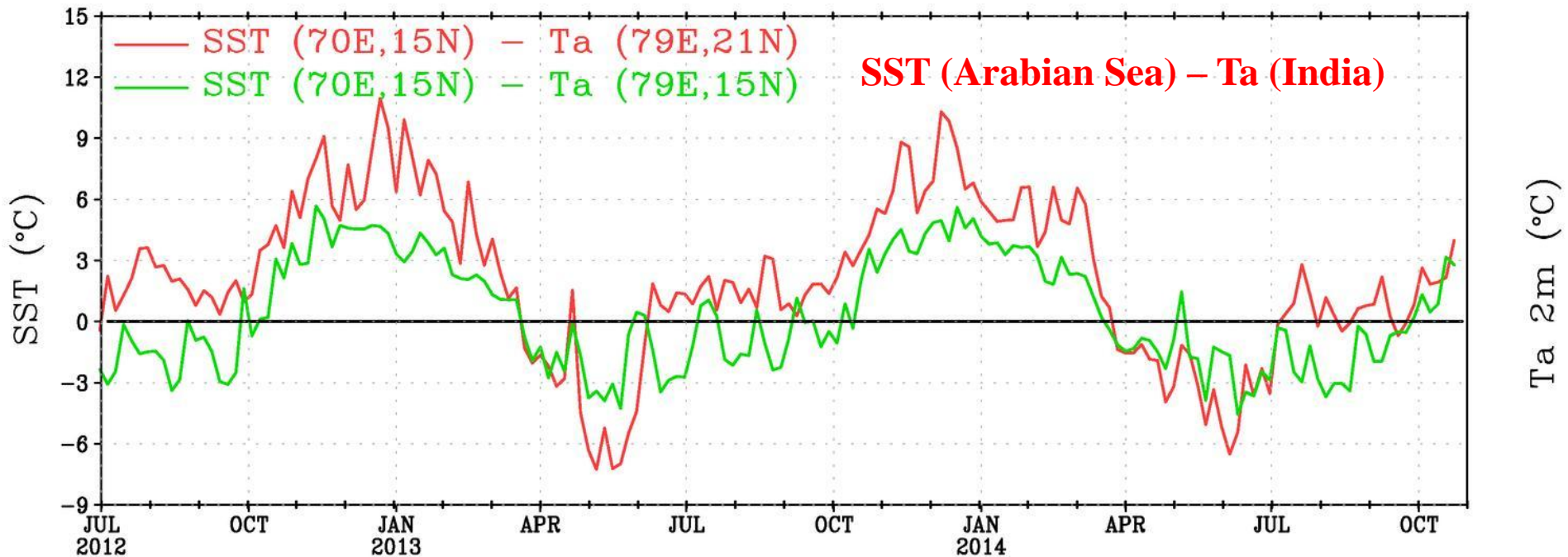
Ta (India)



SST (Arabian Sea)



SST (Bay of Bengal)



Summary

- **The monsoon rain in central Indian Subcontinent is closely linked to the moisture flux from the Arabian Sea, with onset in early June and peak in July/August.**
- **Before the onset of the Indian summer monsoon, the Indian subcontinent is characterized by extremely hot air temperature, with maximum in May. The air temperature immediately cools down with the monsoon onset. There is strong northerly winds at 850 mb in April/May, which coincide with the pre-monsoon drought. Soil moisture shows very dry land before the monsoon onset. The net moisture flux to the Indian subcontinent in May is negative, with weaker flux from the Arabian Sea and stronger out flux into the Bay of Bengal.**
- **Rain along the west coast starts around mid-May, approximately two weeks before rain in the center part of the subcontinent.**
- **The soil moisture increases in sync with the monsoon onset and remains high until the end of October, two months after the peak monsoon rainfall.**
- **Soil moisture measured by AMSR2 is lower in magnitude and range than that by Aquarius in the central India.**

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